

## **CHAPTER 3**

### **BOUNDARY LAYER DYNAMISM AND PM<sub>2.5</sub>-PBLH INTERACTION OVER BRAHMAPUTRA VALLEY AT THE VALLEY-SITE SCALE**

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**3.1 INTRODUCTION**

In the previous chapter, the interaction and spatio-temporal variation in PM<sub>2.5</sub>, AOD 550 nm and meteorological parameters are studied. The concentrations of aerosols are also highly influenced by a change in PBLH. The spatial and diurnal analysis of PBLH and PM<sub>2.5</sub>, and PM<sub>2.5</sub>-PBLH relationship over the complex terrain of BV using data from MERRA-2, ERA5 ECMWF and GPS radiosonde forms the core of investigation in this chapter.

The planetary boundary layer (PBL), the lowest part of the atmosphere, has a profound influence on the atmospheric process, climate system and air quality. It acts as a turbulent buffer (that regulates the exchange of heat, momentum, moisture, dust and aerosols) between the earth's surface and the free atmosphere [1-4]. The exchange of aerosols through the PBL is mostly controlled by the atmospheric processes such as convective turbulence and advection (dispersion, transport, transformation and deposition) [1, 5]. These atmospheric processes in the PBL directly respond to terrain-induced flow modification, heat flux, frictional drag and pollutant emission which are apparent over a complex terrain [1, 2, 6]. Complex terrain not only exerts an important influence on boundary layer flow [6] but also affects the atmospheric transport and mixing of aerosols at a broad range of spatial and temporal scales [7-9]. The vertical scale of this transport and mixing of aerosol with the free atmosphere is determined by PBLH, one of the most relevant boundary layer properties to investigate PBL [9, 10].

PBLH is a fundamental parameter that characterizes the vertical mixing between the earth's surface and free atmosphere through the PBL [3, 4]. The convective and turbulent processes that control vertical mixing within PBLH are complex and vary on a wide range of parameters like space, time scales and terrain. PBLH variability is driven by multiple forcing mechanisms including the diurnal and seasonal cycle of incoming solar radiation, advection, turbulence, synoptic-scale subsidence, and terrain-induced flow modification [1-2, 11-12].

The concentrations of aerosols are highly responsive to changes in PBLH; during day-time high PBLH (convective boundary layers CBL) generated by the convective updrafts, increases vertical diffusion of aerosols and reduces their concentrations. While during night-time low PBLH (shallow stable boundary layer SBL), reduces vertical diffusion of aerosols, and therefore, increases aerosol concentration [13, 14]. In addition, the

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accumulation, dispersion, and fate of aerosol are also dictated by several chemical, physical and dynamic atmospheric processes that are relevant to PBL [5, 11, 15, 16 and the reference therein].

As mentioned earlier, BV is surrounded by mountains and hills in the north, south and east (southeast), and opens towards the Indo-Gangetic Plains (IGP) in the west. The complex terrain of the valley has a cumulative effect on the diurnal and seasonal evolution of PBLH, which in turn plays a significant role in the accumulation of local and long-range transported aerosol and its residence [17]. In addition, the biogenic, carbonaceous and marine sea salt aerosols from the huge forest cover, open burning of solid waste disposal [18] and the Bay of Bengal contribute to aerosols loading over the valley [17, 19-21]. The meteorological conditions associated with the Asian summer monsoon over the Indian subcontinent [22] are implicit in ensuring a typical role in boundary layer dynamics and aerosol environment over the region. Diurnal variability of PBLH is one of the prime factors in the variation of surface PM<sub>2.5</sub>, PM<sub>10</sub>, and black carbon (BC) over the BV [21, 23]. PBLH, in turn, is a function of turbulent mixing, horizontal high wind speeds, fumigation effect and subsidence in PBL over BV [17, 20, 24-26].

The high pollutant load of the IGP enters the BV from the west [21]. Consequently, PM<sub>2.5</sub> concentrations in the largest city, Guwahati, of BV are about five-fold higher than the USEPA standards [17]. The unique topography of the valley, atmospheric processes modified by terrain complexity, boundary layer dynamic, abundance of local sources (particularly biogenic aerosols), entrapment of long-range transported aerosols in the PBL, and the extant state of air quality, make the valley significant to aerosol-PBLH studies. Despite its importance, comprehensive investigation of diurnal changes in PM<sub>2.5</sub> and PBLH and the association is virtually non-existent at the valley-site scale.

This study, therefore, investigates (a) the variability of PBLH and surface PM<sub>2.5</sub> concentration at the spatial and diurnal scale; (b) seasonality of PBLH, PM<sub>2.5</sub> and AOD 550 nm; and (c) the association of PBLH with surface PM<sub>2.5</sub> concentration in different seasons over BV. All these investigations are carried out at the valley-site scale. To the best of the author's knowledge, this study is the first to report the diurnal and seasonal variability of PBLH, and the PM<sub>2.5</sub>-PBLH relationship at the valley-site scale for BV.

In addition, this study also attempts to overcome the limitations of inadequate ground-based data at a high spatial and temporal scales for PBLH and PM<sub>2.5</sub>; a longstanding and major limitation of aerosol-PBLH research over the valley. The data limitations can be effectively overcome by the use of satellite-derived aerosol products and reanalysis data. This study utilizes MODIS onboard Terra and Aqua AOD retrievals, ERA5 reanalysis and MERRA-2 PBLH data, and surface GPS radiosonde PBLH estimates.

### 3.2 DATA, SOURCES AND PRE-PROCESSING

The details of the datasets along with sources used in this study are provided in Table 3.1 and the salient features are discussed in the following sub-sections. The ground-based GPS radiosonde data station is shown in Figure 3.1.

**Table 3.1** Technical specifications of the datasets used.

Parameters	Spatial resolution	Temporal resolution	Data source	Available at
BLH	0.25° x 0.25°	hourly	ERA5 ECMWF	<a href="https://www.ecmwf.int/en/forecasts/datasets/browse-reanalysis-datasets">https://www.ecmwf.int/en/forecasts/datasets/browse-reanalysis-datasets</a>
PBLH	0.5° x 0.625°	hourly	MERRA-2	<a href="https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/">https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/</a>
Temperature, pressure, RH, $\theta$ , $\theta_v$ , WS and WD	point	00:00 h UTC and 12:00 h UTC	GPS radiosonde	<a href="http://weather.uwyo.edu/uppeair/sounding.html">http://weather.uwyo.edu/uppeair/sounding.html</a>
Aerosol components: BC, SO <sub>4</sub> , SS <sub>2.5</sub> , DU <sub>2.5</sub> , OC	0.5° x 0.625°	hourly	MERRA-2	<a href="https://disc.gsfc.nasa.gov/">https://disc.gsfc.nasa.gov/</a>
AOD 550 nm	10 x 10 km	10:30 LST and 13:30 LST	MODIS Terra and Aqua	<a href="https://ladsweb.modaps.eosdis.nasa.gov/search/">https://ladsweb.modaps.eosdis.nasa.gov/search/</a>

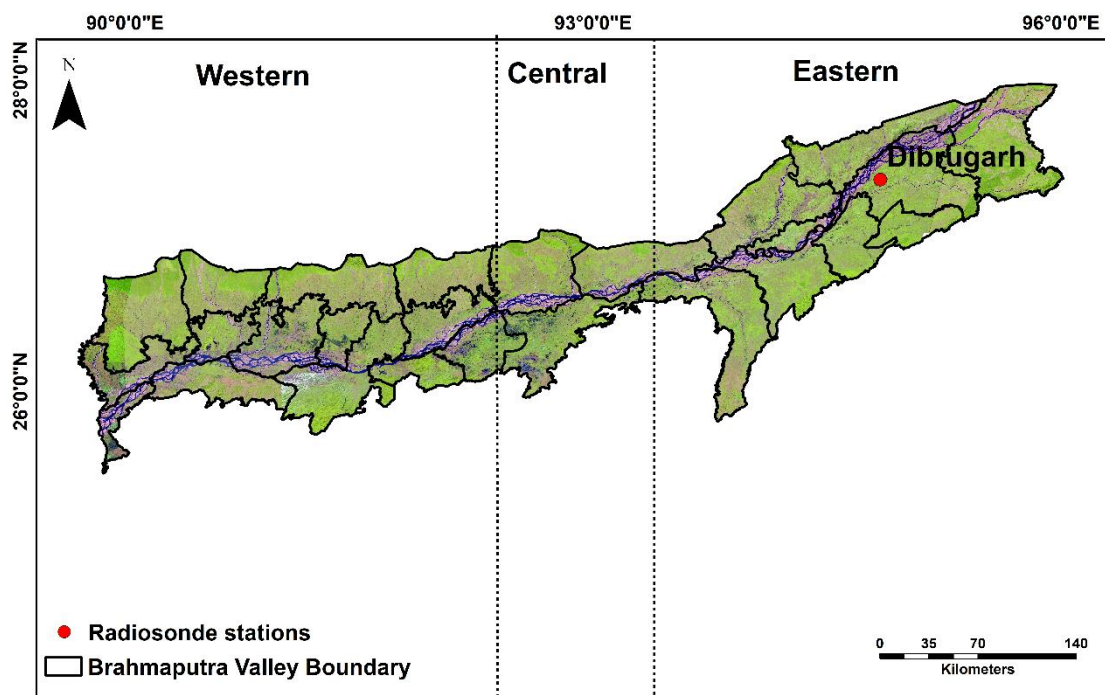
$\theta$ = potential temperature;  $\theta_v$ = virtual potential temperature, WS= wind speed, WD= wind direction, BC= Black carbon, SO<sub>4</sub>= sulfate, SS<sub>2.5</sub>= sea salt, DU<sub>2.5</sub>= dust, OC= organic carbon

#### 3.2.1 ERA5 ECMWF BLH

ERA5 ECMWF reanalysis hourly estimates of boundary layer height (BLH) were downloaded and extracted at 0.25° x 0.25° native spatial resolution for the study period 2016-2020. ERA5 is the fifth-generation global reanalysis product by ECMWF. ERA5

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BLH calculation is based on the bulk Richardson number ( $Ri$ ). BLH is defined as the lowest level at which the bulk  $Ri$  reaches the critical value of 0.25 (where the surface frictional effects are ignored in the computation of the bulk shear due to lack of friction velocity from radiosonde data).



**Figure 3.1** Ground-based station of GPS radiosonde, Dibrugarh.

### 3.2.2 MERRA-2 PBLH

MERRA-2 PBLH hourly time-averaged 2-dimensional data were acquired at  $0.5^\circ \times 0.625^\circ$  native spatial resolution for the study period. This global atmospheric reanalysis was produced by NASA Global Modelling and Assimilation Office (GMAO) using the Goddard Earth Observing System Model (GEOS) version 5.12.4. MERRA-2 PBLH (variable name PBLH) is calculated based on the total eddy diffusion coefficient of heat ( $K_h$ ) with a threshold equal to 10% of the column maximum [27]. PBLH data is an assimilation of modern hyperspectral radiance and microwave observations along with GPS-Radio Occultation datasets and meteorological assimilation (<https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/>).

### **3.2.3 Radiosonde**

GPS-based Radio-sounding system data of upper-air sounding for Dibrugarh station was collected from the Department of Atmospheric Science, University of Wyoming and used for estimation of PBLH. The sounding data are mostly available at the 00:00 h UTC (5:30 h LST) and 12:00 h UTC (17:30 h LST) time scale on daily basis. All estimates of PBLH were done for 00:00 h UTC (5:30 h LST) sounding hours with reference to the height above ground level. The estimated PBLH was used for validation of MERRA-2 and ERA5 reanalysis derived PBLH data for the Dibrugarh station.

### **3.2.4 MODIS AOD 550 nm**

As discussed in section 2.2.1 Chapter 2.

### **3.2.5 PM<sub>2.5</sub> surface concentration**

As discussed in section 2.2.2 Chapter 2.

### **3.2.6 Data pre-processing**

#### **Collocation of data**

For comparison of MERRA-2 PBLH and ERA5 BLH data with reference to radiosonde PBLH estimates, both MERRA-2 PBLH and ERA5 BLH data were extracted at their native spatial resolution  $0.5^{\circ} \times 0.625^{\circ}$  and  $0.25^{\circ} \times 0.25^{\circ}$  respectively. The reanalysis derived PBLH data were collocated in space (geo-referenced taking coordinates of Dibrugarh as the centroid) and time (00:00 h UTC) with reference to radiosonde data. The spatio-temporal collocation of the data was carried out using Python libraries.

#### **Data integration**

MERRA-2 PBLH estimates, MODIS AOD 550 nm retrievals and total PM<sub>2.5</sub> surface concentration (generated from MERRA-2 aerosol components) were re-gridded from native spatial resolution to new target grid size  $0.25^{\circ} \times 0.25^{\circ}$  by using python libraries and IDW (Inverse distance weighting) method on ArcGIS10.5 and interface for BV, for 2016-2020. To examine the diurnal variation of PBLH and the concurrent surface PM<sub>2.5</sub> concentration over the BV, the day-time and night-time mean were used. For seasonality analysis, the seasonal mean of PBLH and (concurrent) PM<sub>2.5</sub> and AOD 550 nm was used.

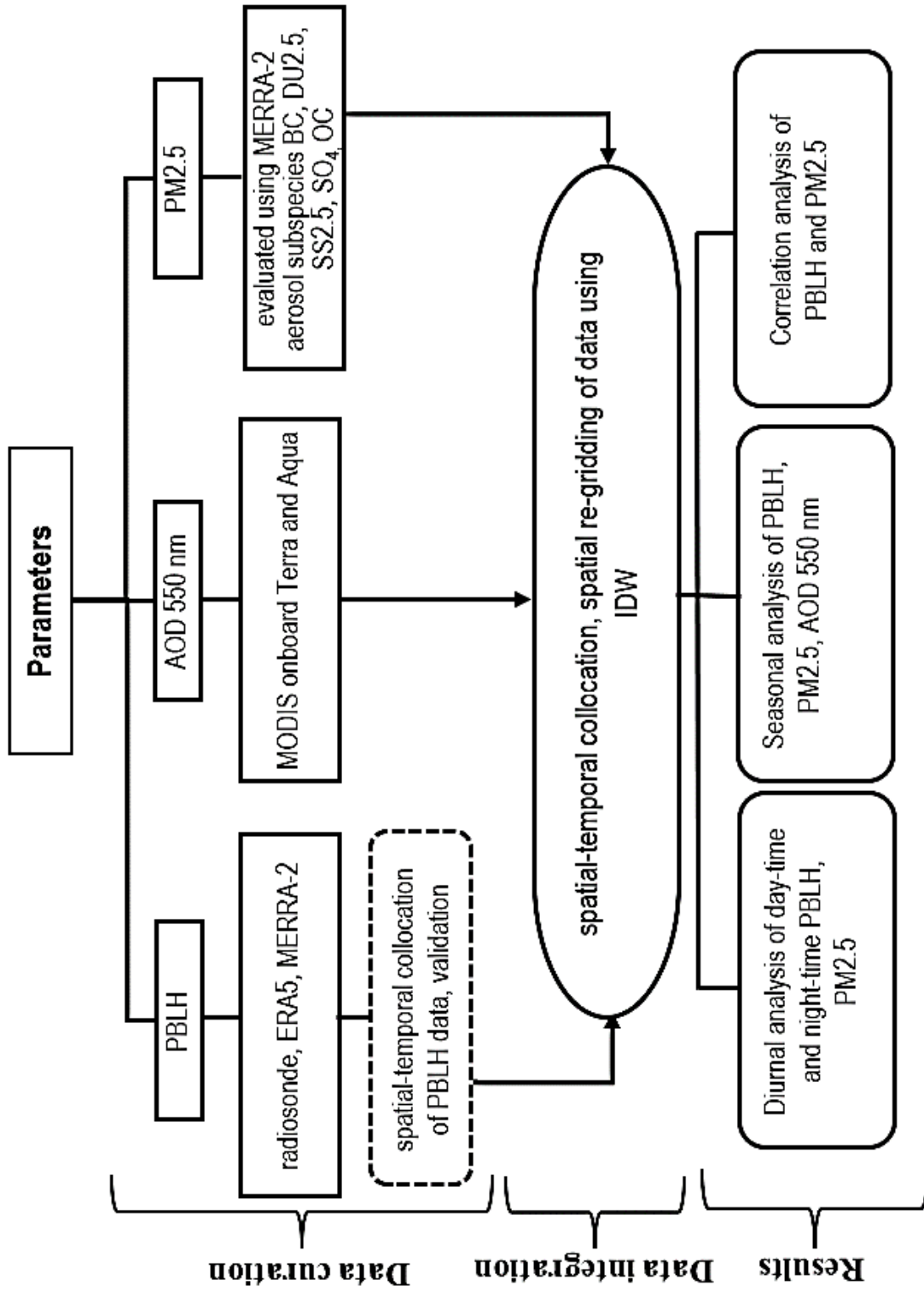


Figure 3.2 Schema of the methodology.

### 3.3 METHODOLOGY

The schema of the methodology used is illustrated in Figure 3.2 and the details are discussed in the following sub-sections.

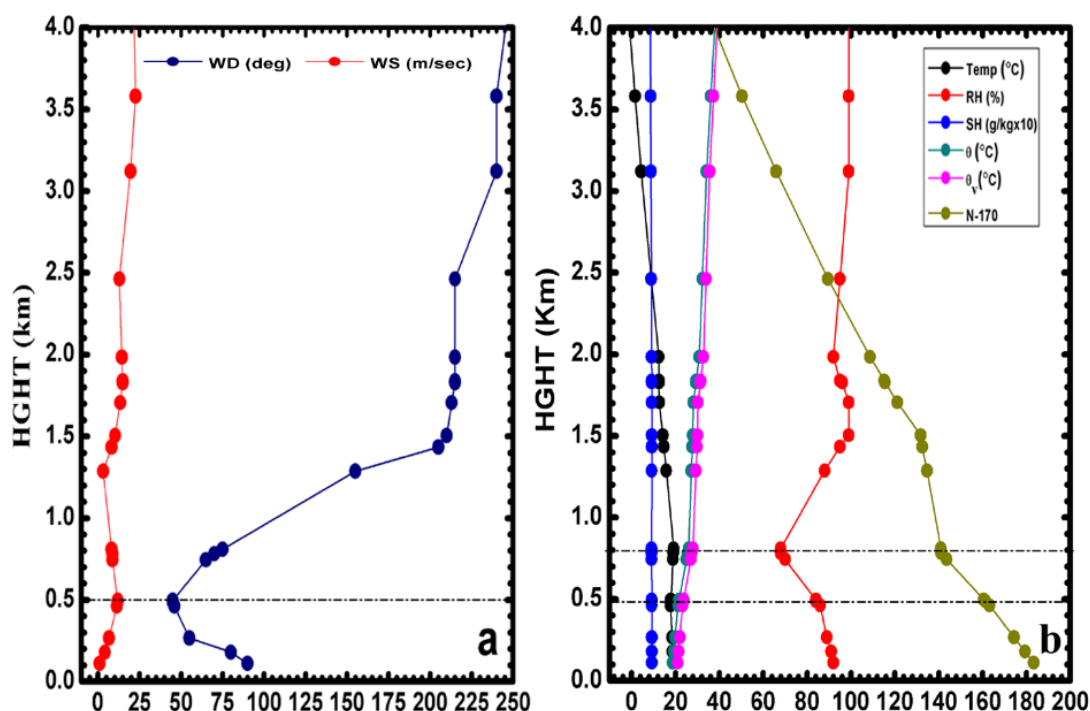
### **3.3.1 Determination of PBLH**

This study investigated the vertical gradient of 8 fundamental atmospheric fields to examine the best indicator of PBLH, over the Dibrugarh station during the early morning hours (00:00 h UTC or 05:30 h LST), 2016 (Figure 3.3a, b). The following 8 vertical gradient-based methods on profiles were examined: (1) maximum height of wind speed, (2) gradient of wind direction, (3) minimum gradient of relative humidity (RH), (4) minimum gradient of specific humidity (q), (5) maximum gradient of temperature ( $^{\circ}\text{C}$ ), (6) maximum gradient of potential temperature ( $\theta$ ), (7) maximum gradient of virtual potential temperature ( $\theta_v$ ), (8) minimum gradient of refractivity (N). Out of all the vertical gradient profiles (Figure 3.3b) generated, wind speed maximum showed the sharpest gradient in the early morning hours (Figure 3.3a). The height of maximum in wind speed profile provides a good measure of mixing height during conditions of atmospheric stability over the site. PBLH, therefore, was determined from the vertical height of the maximum low-level wind speed for 2016-2019.

The atmosphere is considered stable during early morning hours (00:00 h UTC or 5:30 h LST) due to the effect of local topography. For the height of the stably stratified atmospheric boundary layer, characteristic heights of PBLH can be identified from the observed vertical profile of wind velocity [28]. The height of the maximum wind speed vertical gradient provides a good measure of mixing height during the early morning hours, as it is generally well defined from the night-time wind profiles (residual layer) [29] over the land.

Evaluation of PBLH using the Richardson number method was also attempted. However, the lack of wind data (vertical profile) at the same levels as that of the potential temperature and humidity data has led to unreliable results, therefore not included.





**Figure 3.3** Planetary boundary layer height estimates for Dibrugarh station for 00:00 UTC 6 November 2018 using the vertical gradient-based method for eight meteorological parameters. (a) shows vertical profiles of maximum wind speed level and sharp vertical gradient of wind direction, (b) shows vertical profiles of Temperature (Temp °C), potential temperature ( $\theta$ ), virtual potential temperature ( $\theta_v$ ), relative humidity (RH %), specific humidity (SH g/kg x10), and refractivity (N)\*. Estimated PBL heights are shown by dashed horizontal lines. \*Some values of N are shifted for clarity.

### 3.3.2 Validation of PBLH data

For comparison of MERRA-2 PBLH and ERA5 BLH data with reference to radiosonde PBLH estimates, both MERRA-2 PBLH and ERA5 BLH data were extracted at their native spatial resolution  $0.5^0 \times 0.625^0$  and  $0.25^0 \times 0.25^0$  respectively. The reanalysis derived PBLH data were collocated in space (geo-referenced taking coordinates of Dibrugarh as the centroid) and time (00:00 h UTC) with reference to radiosonde data.

ERA5 ECMWF and MERRA-2 derived PBLH was validated with reference to radiosonde PBLH estimates for the Dibrugarh station, for the period 2016 to 2019. This was done based on statistical parameters- coefficient of determination ( $R^2$ ), root mean square error (RMSE), slope and y-intercept (Figure 3.4) (details in section 4.1). MERRA-2 hourly PBLH data were extracted for day-time (10:30 h LST to 01:30 h LST) and night-

time (20:00 h LST to 23:00 h LST) in native spatial resolution, and averaged for diurnal analysis. For seasonal analysis, the mean for different seasons- winter (Jan-Feb), pre-monsoon (Mar-April-May), monsoon (Jun-Jul-Aug-Sept), and post-monsoon (Oct-Nov-Dec) were calculated using monthly PBLH data.

### **3.3.3 Calculation of day-time and night-time PM<sub>2.5</sub> surface concentration**

Day-time and night-time PM<sub>2.5</sub> surface concentration was calculated for BV using hourly data of all the individual major aerosol components BC, DU<sub>2.5</sub>, SS<sub>2.5</sub>, OC and SO<sub>4</sub> employing equation (1), for the period 2016-2020. For diurnal analysis, PM<sub>2.5</sub> surface concentration of day-time (10:30 h LST to 01:30 h LST or 05:00 h UTC to 08:00 h UTC) and night-time (20:00 h LST to 23:00 h LST or 14:30 h UTC to 17:30 h UTC) was calculated and averaged. The monthly mean of PM<sub>2.5</sub> surface concentration was calculated using the monthly mean of all the individual aerosol components using equation (1) (below) given by Hand et al. (2011) [30]. The calculated monthly mean of PM<sub>2.5</sub> surface concentration was averaged for each of the seasons.

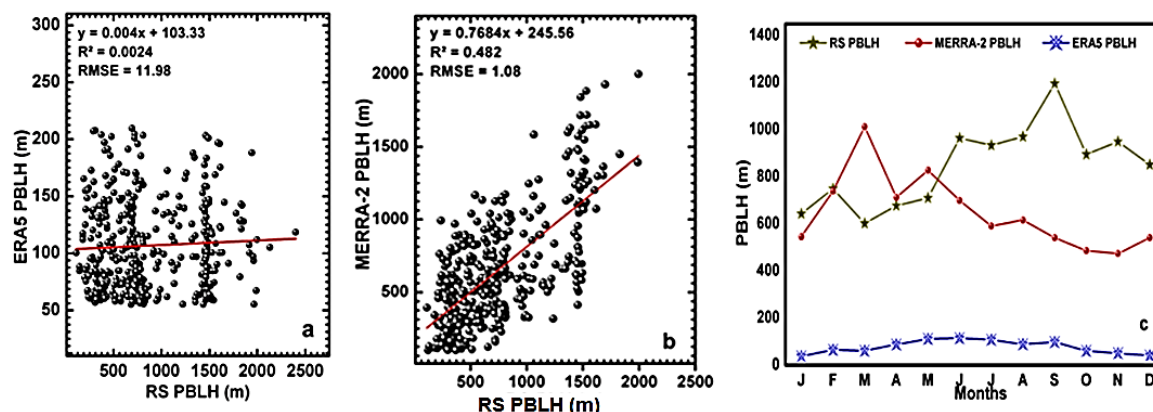
$$\text{PM}_{2.5} = 1.375 \times [\text{SO}_4] + 1.8 \times [\text{OC}] + [\text{BC}] + [\text{DU}_{2.5}] + [\text{SS}_{2.5}] \quad (1)$$

## **3.4 RESULTS AND DISCUSSION**

### **3.4.1 Performance of ERA5 and MERRA-2 PBLH with reference to radiosonde derived PBLH**

The results of performance assessment of MERRA-2 and ERA5 PBLH with respect to radiosonde PBLH estimates showed that MERRA-2 PBLH had better ( $R^2 = 0.482$ ; RMSE = 1.082) represented radiosonde PBLH estimates than that of ERA5 ECMWF ( $R^2 = 0.0024$ ; RMSE = 57.25) for the Dibrugarh station (Figure 3.4). MERRA-2 model PBLH is generated based on the total eddy diffusion coefficient of heat ( $K_h$ ) [27] that represents the turbulent diffusion process; whereas ERA5 BLH is computed based on the bulk Richardson number method that measures local turbulence, and is often unable to properly characterize the turbulent properties of convective boundary layers (stratocumulus, shallow cumulus, and dry convective PBLs) [31]. Since the PBLH estimated in this study was from the vertical gradient of maximum wind speed, PBLH provided by the MERRA-2 dataset had a good agreement with radiosonde estimates. The

performance of MERRA-2 data ( $r = 0.8$ ) is also in line with the finding of Pathak et al. (2019) [32] for the Dibrugarh station.

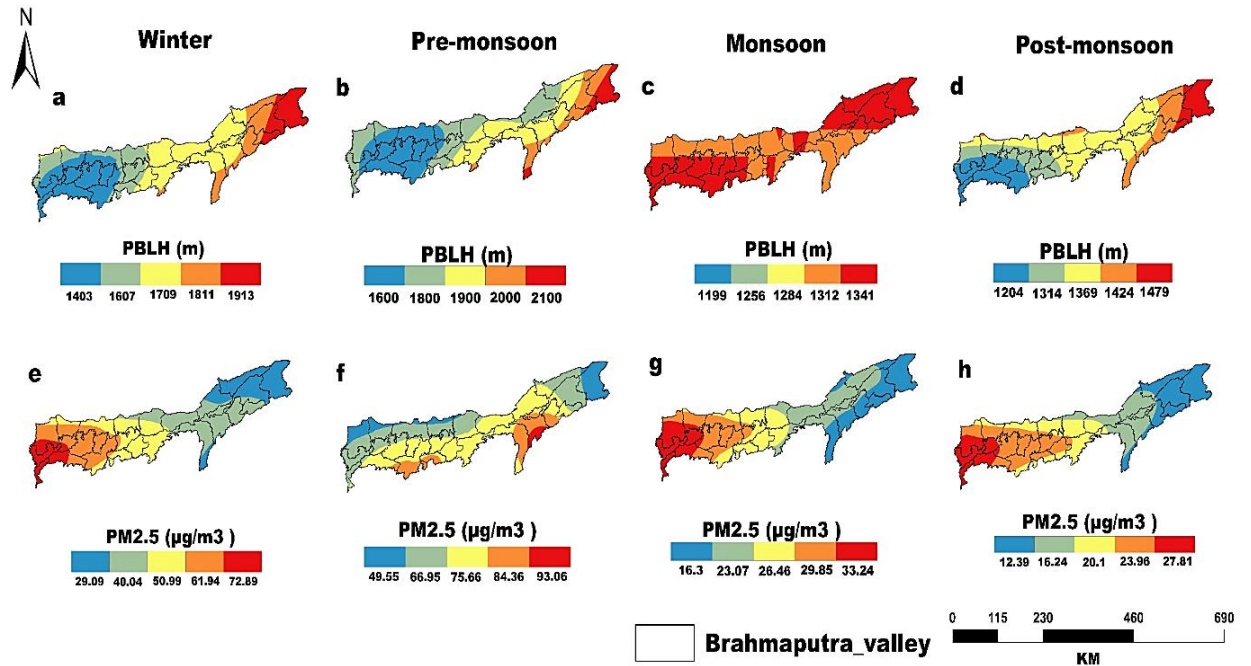


**Figure 3.4** Performance assessment of (a) ERA5 PBLH and MERRA-2 (b) derived PBLH with reference to radiosonde (RS) PBLH estimates of Dibrugarh station at 00:00 h UTC (05:30 h LST), 2016-2019. (c) Comparative analysis of monthly mean of MERRA-2 and ERA5 derived PBLH, and radiosonde PBLH estimates.

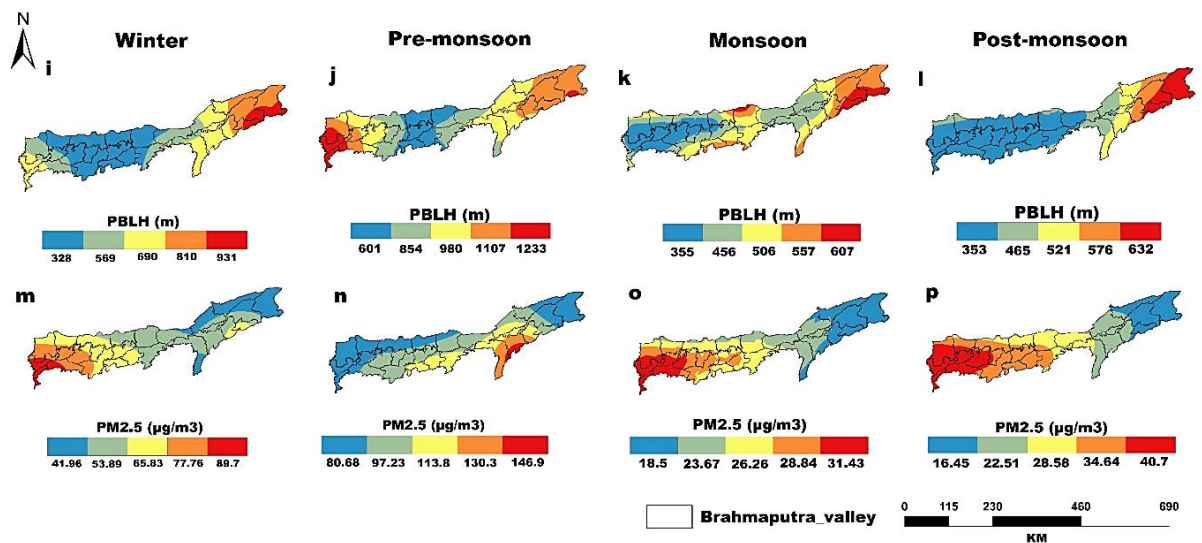
### 3.4.2 Spatial variation of day-time and night-time PBLH and PM<sub>2.5</sub> over BV

High spatial heterogeneity in the diurnal cycle of PBLH structure and PM<sub>2.5</sub> across the valley was observed across seasons. PBLH observed over BV was dominated by strong diurnal variability (Figure 3.5a-d and Figure 3.6i-l) for all the seasons. PBLH was typically shallow (500 m) in evening hours (Figure 3.6i-l), when the surface layer is stable (stable boundary layer) due to infrared radiative cooling. It grows deep (to a couple of kilometres) in the afternoon hours (Figure 3.5a-d) when solar heating causes convective unstable conditions (convective boundary layer). PBLH during nocturnal periods was low, unlike day-hours; attributed to stable conditions with nocturnal inversion- a normal phenomenon over a complex terrain [11].

The deepest day-time PBLH was observed during pre-monsoon (1600-2100 m) closely followed by winter (1400-1900 m), post-monsoon (1200-1470 m) and monsoon (1190-1340 m) over the BV (Figure 3.5a-d). For night-time, shallow PBLH predominates in winter (328-931m) followed by post-monsoon (353-632 m), monsoon (355-607 m), and pre-monsoon (601-1233) m (Figure 3.6i-l).



**Figure 3.5** Spatial variation of (a-d) day-time (10:00 h to 13:00 h LST) mean of PBLH and (e-h) PM<sub>2.5</sub> µg/m<sup>3</sup> surface concentration over Brahmaputra valley in different seasons, 2016-2020.



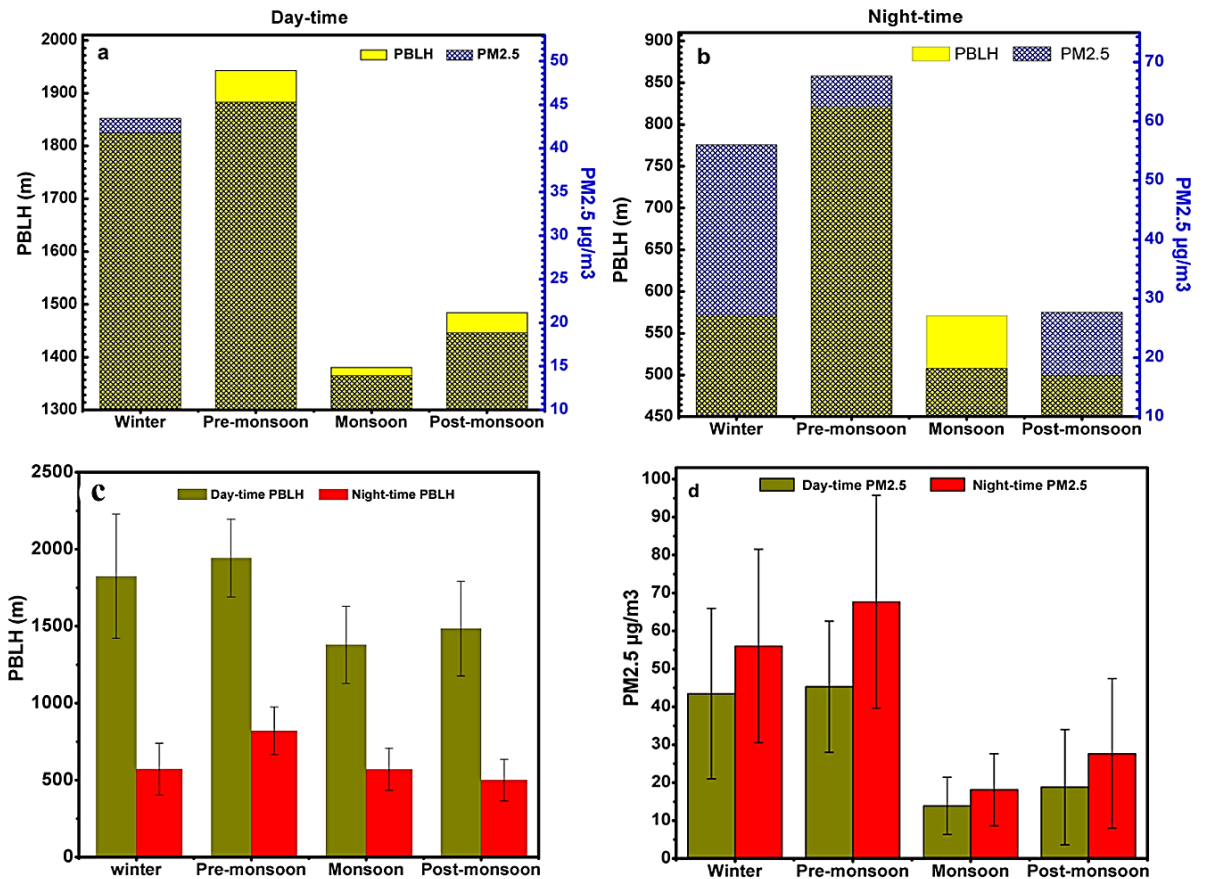
**Figure 3.6** Spatial variation of (i-l) night-time (20:00 to 23:00 h LST) mean of PBLH and (m-p) PM<sub>2.5</sub> µg/m<sup>3</sup> surface concentration over Brahmaputra valley across the seasons, 2016-2020.

For PM<sub>2.5</sub>, a large diurnal variation in PM<sub>2.5</sub> concentration was noted over BV (Figure 3.5e-h and Figure 3.6m-p). Higher PM<sub>2.5</sub> surface concentration was observed at night-time than in the day-time. During day-time PM<sub>2.5</sub> concentration was found to be highest in pre-monsoon (49.55- 93.06  $\mu\text{g}/\text{m}^3$ ), followed by winter (29.09-33.24  $\mu\text{g}/\text{m}^3$ ), monsoon (16.3-33.24  $\mu\text{g}/\text{m}^3$ ) and post-monsoon (12.39-27.81  $\mu\text{g}/\text{m}^3$ ) (Figure 3.5e-h). While for night-time, the highest PM<sub>2.5</sub> concentration was noted for pre-monsoon (80.68- 146.90  $\mu\text{g}/\text{m}^3$ ) followed by winter (41.96- 89.70  $\mu\text{g}/\text{m}^3$ ), post-monsoon (16.45- 40.70  $\mu\text{g}/\text{m}^3$ ) and monsoon (18.5-31.43  $\mu\text{g}/\text{m}^3$ ) (Figure 3.6m-p).

### 3.4.3 Synergism in PM<sub>2.5</sub> and PBLH diurnal variability

Diurnal variation in PM<sub>2.5</sub> concentration in the surface layer was in synergy with the diurnal cycle of PBLH over the BV (Figure 3.7a-d and Table 3.2). The pronounced diurnal evolution of PBLH and PM<sub>2.5</sub> over the valley, reveals that turbulent mixing of air (mechanical turbulence generated by wind shear) plays, understandably, a dominant role in the entrainment of aerosols in PBL. Day-time PBLH (Figure 3.7a) was much higher than that of night-time (Figure 3.7b), while PM<sub>2.5</sub> concentration displays the reverse trend of being much lower during day-time (Figure 3.7a) than at night-time (Figure 3.7b). Lower day-time PM<sub>2.5</sub> concentration was possibly the outcome of vertical mixing caused by thermal and mechanical turbulence within the PBLH. The diurnal heating and cooling of hill slopes often develop thermal turbulence over complex topography [11]. It enables lifting dust and other aerosol particles from the surface and spreading them throughout the entrainment zone of PBLH. During the night, nocturnal radiative cooling of the surface leads to the formation of a ground-based stable inversion layer (nocturnal inversion), that confines aerosol mixing to a shallow layer near the surface, increasing their surface concentration.

PBLH generally becomes shallower under the influence of large-scale subsidence (downward motion), which can weaken the vertical mixing of aerosols and increase surface accumulation [33] over the area.



**Figure 3.7** Comparative analysis of the (valley-site scale) mean (a) day-time PBLH vs day-time PM2.5, (b) night-time PBLH vs night-time PM2.5, (c) day-time vs night-time PBLH, (d) day-time vs night-time PM2.5 across the seasons, Brahmaputra valley, 2016-2020. Vertical error bar represents the standard deviation in dataset.



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**Table 3.2** Descriptive statistics of daytime and night-time PBLH and PM<sub>2.5</sub> in different seasons, Brahmaputra valley, 2016-2020.

Seasons	Statistics	Day-time		Night-time	
		PM <sub>2.5</sub> ( $\mu\text{g}/\text{m}^3$ )	PBLH (m)	PM <sub>2.5</sub> ( $\mu\text{g}/\text{m}^3$ )	PBLH (m)
Winter	Mean	50.38	1580.37	62.83	510.35
	Min	30.70	1238.91	32.70	321.11
	Max	77.56	1865.24	93.70	827.03
	Median	46.94	1579.30	61.95	506.60
	SD	12.09	167.68	14.68	151.12
Pre-monsoon	Mean	49.81	1797.21	70.42	866.34
	Min	40.84	1491.89	51.02	586.44
	Max	57.18	2060.13	101.98	1234.06
	Median	51.71	1765.11	68.83	851.85
	SD	4.72	148.43	12.58	163.63
Monsoon	Mean	16.20	1219.39	21.14	478.27
	Min	11.50	1007.97	13.66	372.64
	Max	23.54	1498.15	31.66	629.59
	Median	15.44	1221.27	19.74	472.21
	SD	3.48	139.10	5.47	77.71
Post monsoon	Mean	20.32	1292.27	30.36	422.44
	Min	8.92	1043.23	11.29	259.90
	Max	43.38	1535.25	59.77	702.46
	Median	17.99	1297.91	27.54	409.17
	SD	8.54	133.17	11.85	97.78

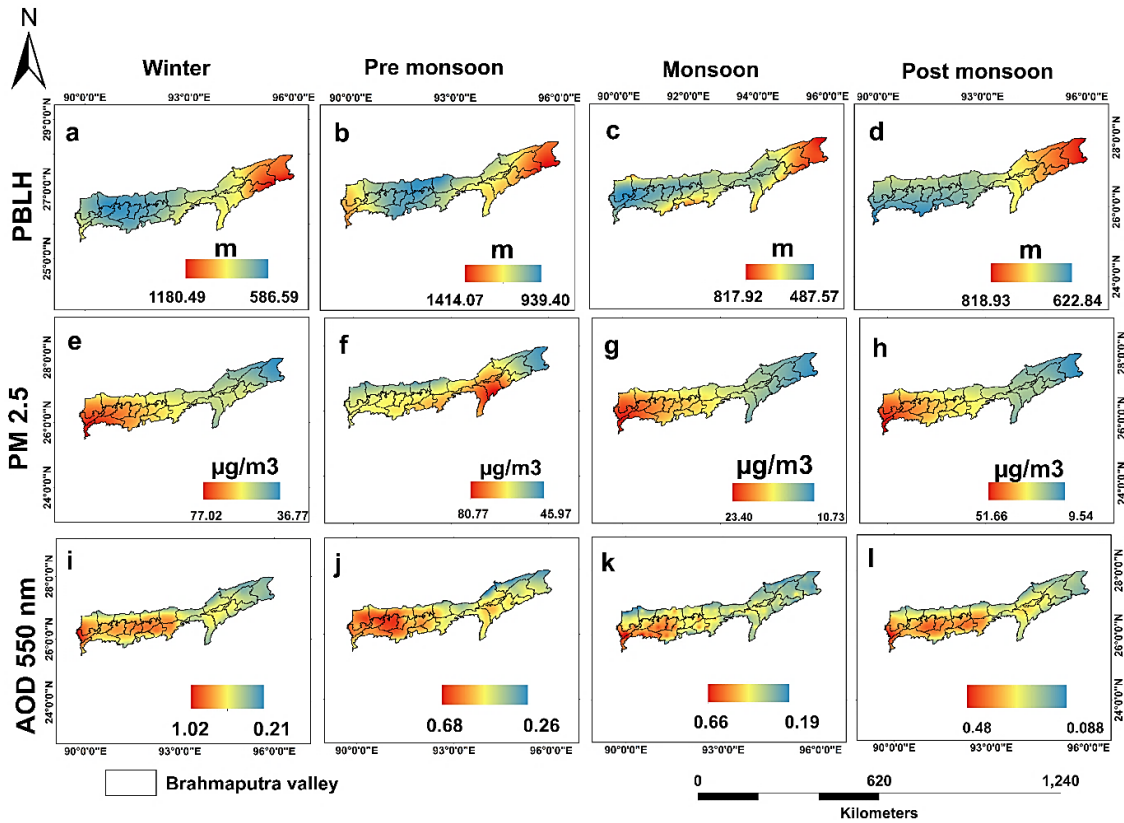
PBLH= Planetary boundary layer height

### 3.4.4 Seasonal changes of PM<sub>2.5</sub>, AOD 550 nm and PBLH over BV

The climatology of the PBLH, especially its seasonal variability, is important for air pollution related studies. PBLH derived from MERRA-2 was utilised to represent the seasonality of PBLH across the spatial domain of BV. The seasonal mean of PBLH, AOD 550 nm and surface PM<sub>2.5</sub> concentration of the entire BV are represented in Figure 3.8. PBLH was observed higher in pre-monsoon closely followed by winter, post-monsoon and monsoon. PBLH was the deepest in pre-monsoon (939.40 to 1414.07 m), and the shallowest in monsoon (487.57- 817.92) over the BV (Figure 3.8a-d).

For all the seasons, PM<sub>2.5</sub> surface concentration and AOD 550 nm displayed high spatial variability similar to PBLH variation. This pattern is in conformity with observations from other research work [34-35]. Relatively larger amounts of PBLH with lower levels of PM<sub>2.5</sub> (Figure 3.8e-h) and AOD 550 nm (Figure 3.8i-l) prevailed over the eastern

parts of BV, compared to the western and central. Deeper PBLH over the eastern BV could be due to the higher turbulence developed over the rough landscape (of the eastern BV), relative to the turbulence developed in the other parts. The eastern part of BV, which comprises hill ranges, dense vegetation cover and vast floodplains, is suitable for developing strong mechanical and convective turbulence (eddy motion). The results indicate that turbulence within the PBLH is effective in regulating the PM<sub>2.5</sub> surface concentration across the entire valley.

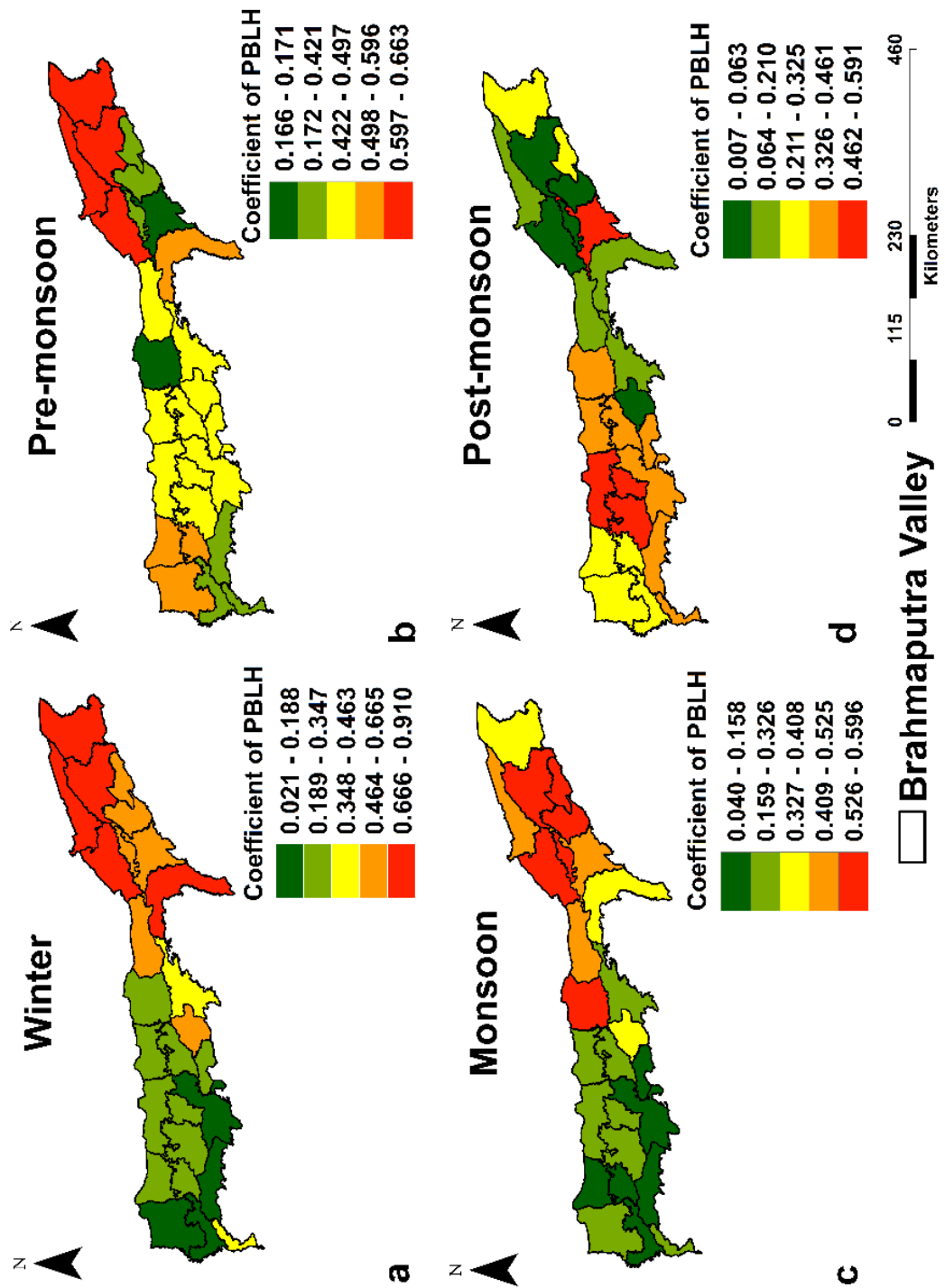


**Figure 3.8** Variation in seasonal mean of (a-d) PBLH, (e-h) PM<sub>2.5</sub> concentration and (i-l) AOD 550 nm over Brahmaputra valley, 2016-2020.

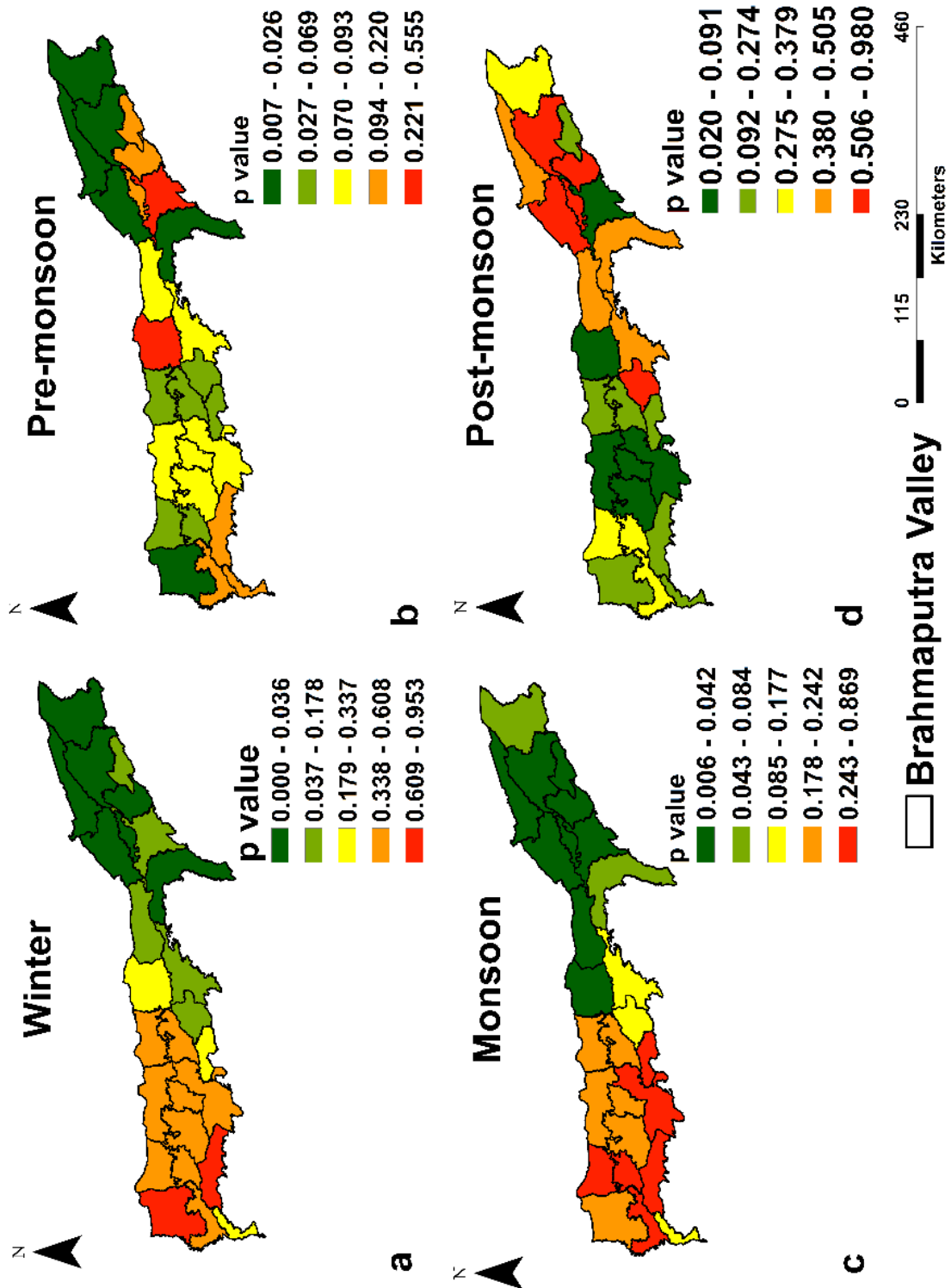


### 3.4.5 PBLH-PM2.5 relationship

Positive association between PBLH and PM2.5 was observed for all the seasons (Figure 3.9a-d). In winter, the eastern parts of BV showed relatively stronger positive relation ( $r=0.66$  to  $0.91$ ;  $0.0001 \leq p \leq 0.17$ ), than the central ( $r=0.34$  to  $0.66$ ;  $0.17 \leq p \leq 0.33$ ) and western parts ( $r=0.02$  to  $0.46$ ;  $0.33 \leq p \leq 0.95$ ) (Figure 3.9a and Figure 3.10a). During pre-monsoon, most areas in the central and western parts of the valley showed weak to moderate positive relation ( $r=0.172$  to  $0.59$ ,  $0.02 \leq p \leq 0.22$ ), whereas eastern parts showed strong positive relationship ( $r=0.59$  to  $0.66$ ;  $0.007 \leq p \leq 0.02$ ) (Figure 3.9b and Figure 3.10b). In the monsoon, weak ( $r=0.04$  to  $0.32$ ;  $0.08 \leq p \leq 0.86$ ) to moderate ( $r=0.40$  to  $0.59$ ;  $0.006 \leq p \leq 0.008$ ) relationship predominated over the valley (Figure 3.9c and Figure 3.10c). Whereas, poor to strong PBLH-PM2.5 relationship ( $r=0.007$  to  $0.59$ ;  $0.02 \leq p \leq 0.98$ ) prevailed in the post-monsoon season (Figure 3.9d and Figure 3.10d). The eastern BV represented unusual strong positive PBLH-PM2.5 relationship consistently for the three consecutive seasons: winter, pre-monsoon and monsoon. Eastern BV is less polluted and represents the highly heterogeneous terrain, and forested landscape as compared to the western and central parts. Low p-value [ $\leq 0.05$  correlation significant at the 0.05 level (2-tailed)] suggests that changes in the PM2.5 are associated with changes in PBLH.



**Figure 3.9** Spatial distribution of Pearson's correlation coefficient between PBLH and  $\text{PM}_{2.5}$   $\mu\text{g}/\text{m}^3$  over Brahmaputra valley, 2016-2020, in (a) winter, (b) pre-monsoon, (c) monsoon and (d) post-monsoon.



**Figure 3.10** Spatial distribution of p-values of the correlation coefficient between PBLH and PM2.5  $\mu\text{g}/\text{m}^3$  over Brahmaputra valley, 2016-2020, in (a) winter, (b) pre-monsoon, (c) monsoon, and (d) post-monsoon. p-values  $\leq 0.05$  indicate significance at the 0.05 level (2-tailed).

### **3.5 CONCLUSION**

Diurnal variation of PBLH and PM<sub>2.5</sub> was pronounced. Deeper PBLH was observed during the day-time. The diurnal variability of PM<sub>2.5</sub> surface concentration was in synergy with the diurnal variation of PBLH over BV. During day-time deeper PBLH (up to 2100 m in pre-monsoon) contributed to low PM<sub>2.5</sub> surface concentration due to convective turbulent mixing developed over the valley. While at night-time nocturnal radiative cooling of the surface caused shallow PBLH confining aerosol vertical mixing thus increasing surface concentration of PM<sub>2.5</sub>. The pronounced diurnal variability of PBLH over BV is possibly driven by terrain-induced flow modification due to the heating and cooling of hill slopes over BV, a common phenomenon over complex terrain. Diurnal variation in PM<sub>2.5</sub> surface concentration showed strong synergy with the diurnal variation of PBLH over the BV. PBLH was observed to be much higher in the day-time than at night-time, and PM<sub>2.5</sub> concentration was lower during the day-time than night-time.

A clear synergy existed between PBLH and PM<sub>2.5</sub> surface concentration in all the seasons over the valley. Shallow to deep PBLH was observed from west to east BV, which could be due to the localised topographic influence on the turbulence process. The surface heterogeneity of the eastern part is more as compared to other parts. Dense vegetation cover, small hill ranges and vast floodplains prevailing over the eastern zone could favour the development of strong mechanical and convective turbulence (eddy motion). The findings reveal that the PBLH structure varied with the underlying surface roughness and the turbulence generated across the valley, which in consequence affects the surface concentration of PM<sub>2.5</sub> and AOD 550 nm.

PBLH-PM<sub>2.5</sub> relationship showed remarkable spatial variation over the valley. The easternmost parts of the valley had consistently represented a strong positive PBLH-PM<sub>2.5</sub> relationship for three consecutive seasons- winter, pre-monsoon and monsoon; while the central and western parts were noted with weak to moderate positive relation. The findings of this study are crucial for understanding the aerosols-PBLH feedback mechanism, and could be effectively used to determine the regulating factors of air quality at the regional level.

This study also explored different data sources for PBLH- an important step for a region otherwise known for data scarcity. MERRA-2 PBLH outperformed ERA5 ECMWF BLH

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when compared with radiosonde PBLH estimates for the Dibrugarh station, for 2016-2019. Since MERRA-2 PBLH is retrieved based on the total eddy diffusion coefficient of heat ( $K_h$ ) that captures the vertical turbulent diffusivity, it represents the radiosonde PBLH estimates. This study also shows that the vertical gradient of WS is an effective method for estimating PBLH for early morning hours over a complex terrain.

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